

FINDHORN TERRACES

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OS Grid Reference: NH845366

Highlights

This site demonstrates a particularly good assemblage of glacial outwash and river terraces formed respectively during and following the melting of the Late Devensian ice-sheet.

Introduction

The site (NH 845366) is located on the southern side of the middle River Findhorn, within the Streens Gorge, 20 km south of Nairn near the settlement of Ballachrochin. It is notable for a series of glaciofluvial and fluvial terraces (Figure 7.16), which occupy the lower part of the north-west facing slope of Carn Torr Mheadhoin (543 m OD) and are cut into the extensive glacial and glaciofluvial deposits found throughout the Streens Gorge (Young, 1980). The area is described by Horne (1923), Young (1980) and Auton (1990b).

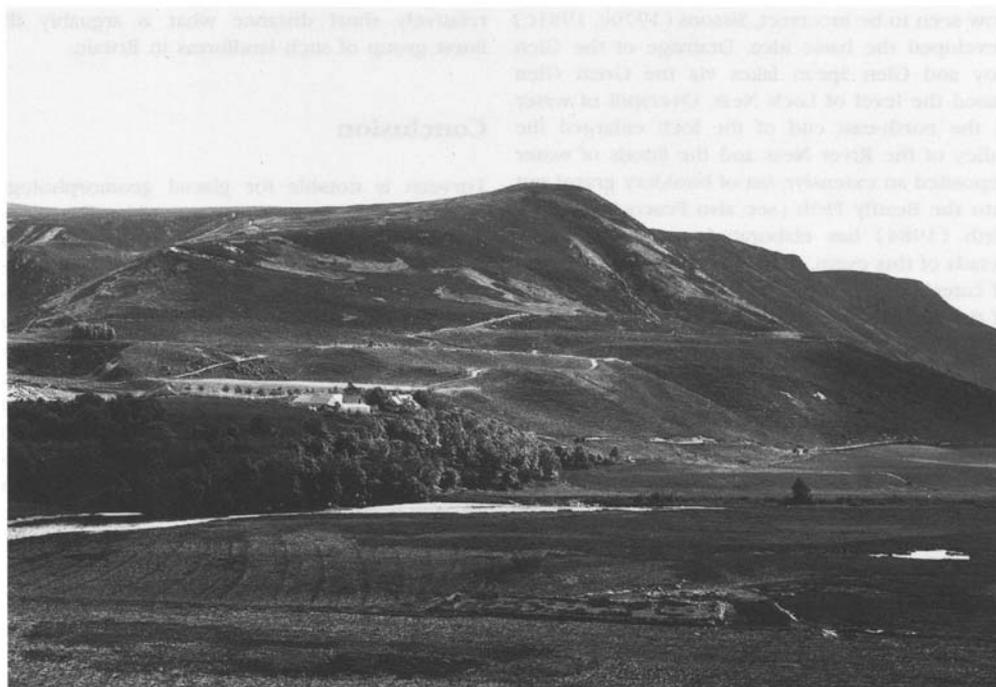


Figure 7.16: The Findhorn terraces at Ballachrochin. (British Geological Survey photograph C1415.)

Description

Horne originally identified eleven terrace levels. More recent mapping by Auton (1990b) has shown that there are thirteen terraces, of which the lowest five occur at 245–275 m OD and exhibit downvalley gradients of 35–50 m per km⁻¹. These terraces locally abut terrace 6 at 285 m OD. In section this flat-topped feature comprises 1.0 m of clast-supported, well-rounded gravel underlain by 1.5 m of a horizontally laminated, low-angle cross-bedded, silty, fine-grained sand. This sand in turn passes down into 2.0 m of finely interlaminated sandy silt and clay with dropstone cobbles and sparse interbeds of diamicton. Above terrace 6, terraces 7 to 11 extend from 287–310 m OD, with the terrace at 305 m OD containing a small steep-sided circular kettle hole 5 m deep. By contrast, terrace 12 (at 340 m OD) is cut into bedrock. The sequence ends at 365 m OD with a small outwash fan on the western side of the Allt a'Choire Bhuidhe.

Interpretation

Horne (1923) interpreted the terraces as fluvial features, although accepting that some of the higher levels were probably glaciofluvial in origin. Young (1980) regarded them as eskers. Auton (1990b), in the most recent investigation, interpreted the landforms as kame terraces and thus of glaciofluvial origin, being closely related to the downwasting of an isolated mass of stagnant ice.

A key part of the sequence in Auton's interpretation of the site is terrace 6 (at 285 m OD). This he considers to be the remains of a glaciolacustrine delta, since the sedimentary sequence closely resembles that of the lower part of the Malaspina delta in Alaska, as described by Gustavson *et al.* (1975). Such an interpretation is not new, having already been anticipated in part by Horne (1923). However, this reconstruction clearly requires the presence of a temporary glacial lake. Young (1980) claimed that the higher terraces are eskers and as such do not require the existence of a glacial lake within the valley, as suggested earlier by Bremner (1939c) and Charlesworth (1956). Auton rejected Young's interpretation and developed a model in which most of the landforms in this middle part of the Findhorn Valley are of paraglacial origin, that is they were formed by 'non-glacial processes that are conditioned by glaciation' (Church and Ryder, 1972). In particular, he considered that the terrace sequence at Ballachrochin developed in response to a stagnating ice mass in the Streens Gorge, which steadily downwasted during the Late Devensian and in so doing created local, temporary glacial lakes. Successive ice margins have been reconstructed by Auton at 460 m, 400 m, 380–350 m, 340–300 m (310 and 305 m benches cut at this stage), 300–260 m (benches between 255 and 287 m cut at this stage) and 250 m OD (final benches cut after this stage).

All the major river valleys in upland Scotland possess sets of terraces which are of fluvial and glaciofluvial origin. It is unusual, however, to find staircases of terraces which extend 80 m above the valley floor and possess 13 identifiable benches. This site on the River Findhorn is notable on both accounts.

The flight of terraces is one of the highest and most remarkable in Scotland, the sequence of 13 levels being related to a complex pattern of deglaciation in this part of the middle Findhorn Valley. Although the site has recently been investigated in considerable detail in terms of its glacial history, the Holocene development of the lower, fluvial, terraces has yet to be attempted. Only when this has been completed will the full significance of the site be disclosed.

Conclusions

The principal landforms at this site comprise a sequence of glacial outwash and river terraces. They are remarkable for the number of levels present and their altitudinal extent. Their development reflects the complex pattern of melting and wastage (deglaciation) of the last (Late Devensian) ice-sheet in the area (approximately 14,000–13,000 years ago). The site represents a striking example of terraces formed by glacial meltwater and river processes during and following deglaciation.

Reference list

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